

# Model based distributed water balance monitoring of the White Volta catchment in West Africa through coupled meteorological-hydrological simulations

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**Abstract.** Sustainable water management requires the quantification of the spatial and temporal changes of water balance variables. Fully distributed hydrological simulations of these variables are especially in regions with weak infrastructure challenging, because the required meteorological input data are often not available in a sufficient spatial and temporal resolution. One possibility to deal with this limitation is to provide the required input data with a meteorological model. This combination results in a one way meteorological-hydrological coupling system. Within the framework of the GLOWA-Volta project it is investigated to what extent meteorological models are able to provide the required meteorological fields with sufficient accuracy for the hydrological modeling. For this study the mesoscale meteorological model MM5 and the fully distributed water balance simulation model WaSiM-ETH were first adapted and validated separately. The research area is the White Volta catchment in the semi-arid to sub-humid climate zone in West Africa. The meteorological simulations tend toward overestimating measured precipitation sums. The coupled meteorological-hydrological runoff simulations show similar model performances as the simulations driven by observations indicating the potential of this system for a contemporary estimation of the terrestrial water balance.

changes of water fluxes in the catchment. This information is of central importance for decision making in water resources management in the Volta basin.

In the northern part of the basin rain-fed agriculture is the major source of livelihood. Due to high population growth the demand for water supply, food and energy production increases permanently. The energy production, especially in Ghana, which shares around 40% of the Volta basin, is directly linked to water availability because the main energy source is hydropower at Akosombo, a dam which was built in 1965 and impounds since then Lake Volta, one of the largest artificial lakes in the world. Because of weak infrastructure in the basin, observations of meteorological and hydrological variables are rare and no basin-wide long-term time series are available. Furthermore, there is a delay of more than one year before measuring data of the complete basin are collected and digitised. This means that for near-term decision making the required information is not available by measuring data. Thus, it is investigated if a model based approach may provide this information. For the model based approach the results of the meteorological simulations are used as input data for the hydrological modeling.

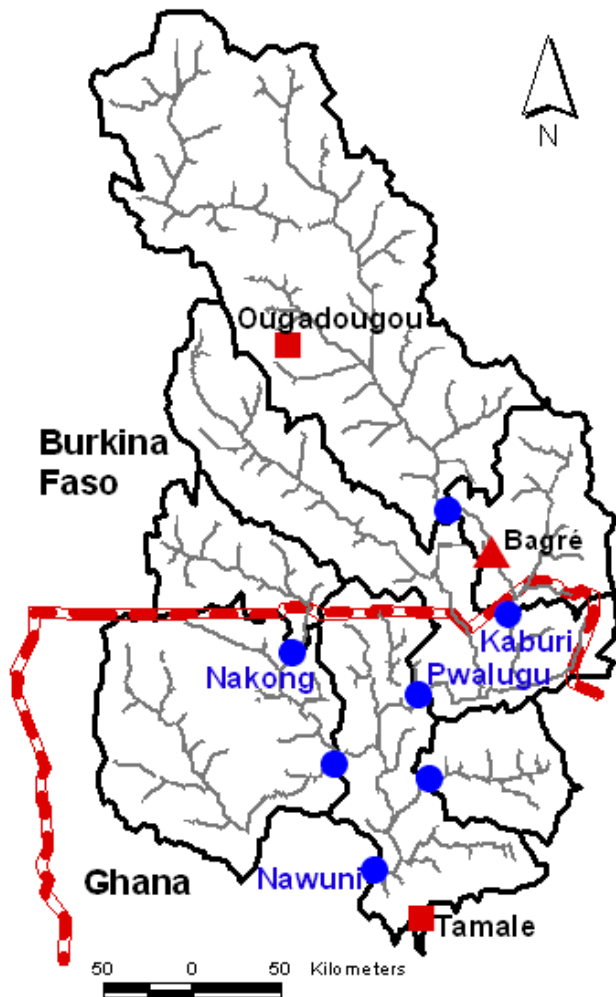
## 1 Introduction

Sustainable water management in the Volta Basin of West Africa is hampered by the fact that only little hydro-meteorological information is available. The central question to be solved therefore is, whether a mainly model based water balance monitoring system can be used to provide scientifically sound quantification of the spatial and temporal

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## 2 Research area

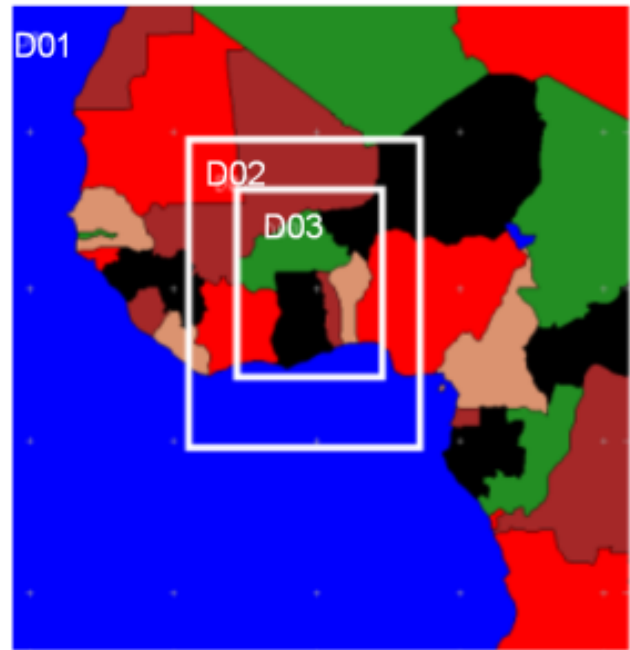
This one-way coupling system was applied for one of the main tributaries of the Volta basin, the White Volta catchment (94 044 km<sup>2</sup>) which is situated upstream of Lake Volta in Northern Ghana and Burkina Faso (Fig. 1). The White Volta basin is very flat particularly in the southern part (<0.1%). The predominant land use types are Guinea savannah in the southern and Sudan savannah in the northern part. The main geological systems of the basin are a Precambrian platform and a sedimentary layer, the Voltaian sandstone basin. The predominant soil types are lixisols in the southern and arenosols in the northern part (Jung, 2006). Since 1993 the



**Fig. 1.** Location of subcatchments and their outlets in the White Volta catchment (resolution: 1 km).

White Volta catchment is disturbed by a dam and hydropower generation in Bagré in the South of Burkina Faso. For this reason and data availability the study focuses on the Ghanaian part of the White Volta basin.

The Volta basin is situated in the semi-arid to sub-humid climate zone with mean annual temperature between 27 and 36°C in the northern and between 24 and 30°C in the southern part. Mean annual precipitation ranges from less than 300 mm (North) to more than 1500 mm in the South whereof around 80% falls between July and September. Evapotranspiration is a very important factor in this region. The mean annual potential evaporation lies between 2500 mm in the North and 1500 mm in the South. Approximately 80% of the precipitation amount evapotranspires during the rainy season (Oguntunde, 2004).

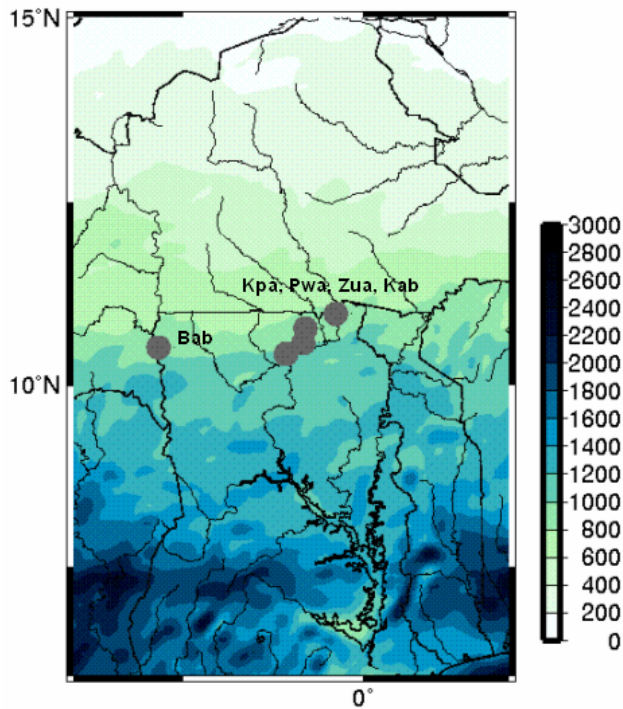


**Fig. 2.** Nesting strategy (3 domains) for the meteorological modeling with MM5 for the Volta basin.

### 3 Meteorological modeling

For the meteorological simulations the mesoscale meteorological model MM5 (Grell et al., 1994) was used. For this study MM5 was applied in non-hydrostatic mode to dynamically downscale the global atmospheric fields from 1° to 9 km resolution stepwise using three domains with horizontal resolutions of 81×81 km<sup>2</sup> (61×61 gridpoints), 27×27 km<sup>2</sup> (85×67 gridpoints) and 9×9 km<sup>2</sup> (157×121 gridpoints) (Fig. 2). The required global analysis fields are obtained from the National Centers for Environmental Prediction (NCEP) with a temporal resolution of 6 h. For the vertical resolution 26 layers from the surface up to 30 hPa were chosen. In Domain1 available observations obtained from radiosondes were incorporated into the simulations to include vertical profiles of atmospheric variables into the modeling process. Kunstmann and Jung (2003) determined an adequate configuration of the available parameterisations for the Volta basin which was used for this study. These are the OSU-Land-Surface Model (Chen and Dudhia, 2001), the MRF-PBL scheme (Hong and Pan, 1996) for the planetary boundary layer, the convective (i.e. cumulus) parameterisation according to Grell et al. (1994), the microphysics according to Reisner et al. (1998) (Mixed phase Graupel) and the cloud-radiation scheme (Grell et al., 1994). For this study the MM5 version 3.6 in the one-way nesting approach was used. The output was saved every 3 h.

The simulation period was the year 2004. Figure 3 shows the spatial distribution of the simulated annual precipitation

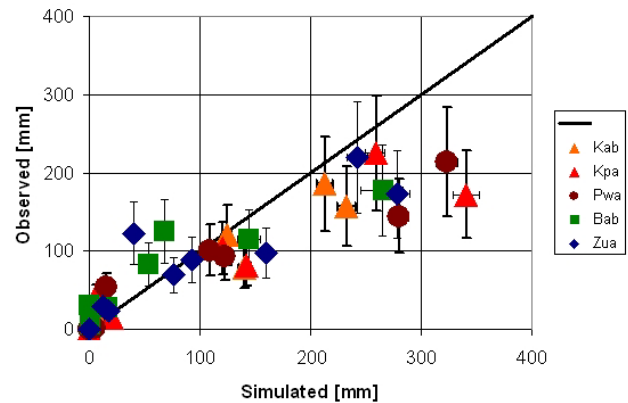


**Fig. 3.** Simulated annual precipitation sum for 2004 (Domain3:  $9 \times 9 \text{ km}^2$ ) for the Volta basin and location of observation stations Kaburi, Kpasenkpe, Pwalugu, Babile and Zuarungu (North Ghana).

[mm] for Domain3 ( $9 \times 9 \text{ km}^2$ ). The precipitation distribution shows a strong North-South gradient with values of less than 300 mm in the North and over 1800 mm in the South. The scatter plot in Fig. 4 shows a comparison of the available observed and simulated monthly precipitation sums, because so far no continuous time series were available for the complete year 2004. Additionally, the scatter plot shows the variability of both the observation and the simulation. The coefficient of variation of the intra  $9 \times 9 \text{ km}^2$ -scale rainfall variability ranges between 0.25 and 0.4 (Friesen, 2003). Figure 4 shows a fairly well agreement for low and medium rainfall sums but for months with heavy rainfall the simulation overestimates the measured rainfall amount significantly. The coefficient of determination  $R^2$  varies between 0.80 and 0.91. For the comparison it has to be considered that a point measurement (observation) is compared here with a result of the meteorological model which is only able to simulate precipitation averages on a scale that is 2–4 times the model resolution. This is especially relevant for regions with a high spatial variability.

#### 4 Hydrological modeling

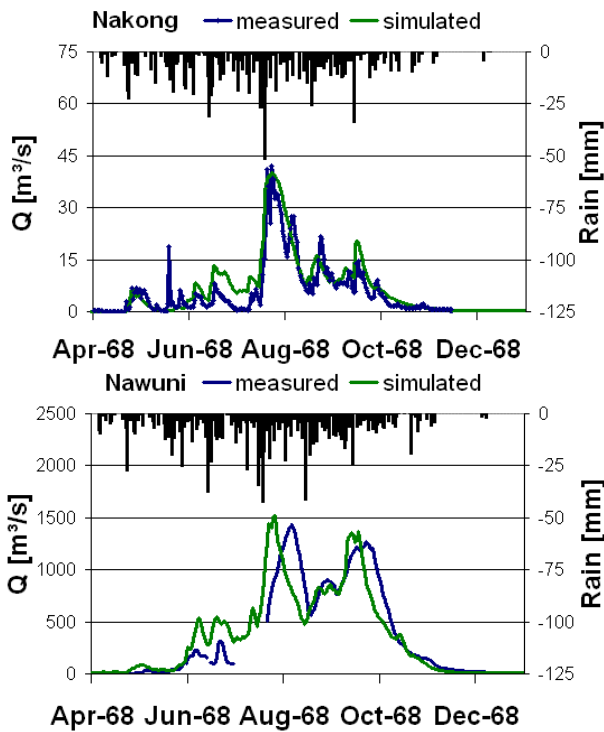
For the hydrological simulations the Water balance Simulation Model WaSiM-ETH (Schulla and Jasper, 2000) was used. It is a deterministic, fully distributed modular model



**Fig. 4.** Comparison of simulated and measured monthly precipitation sums in 2004 for Kaburi, Kpasenkpe, Pwalugu, Babile, Zuarungu.

for the simulation of the hydrologically important parts of the water balance and uses physically based algorithms for most processes. Fluxes in the unsaturated zone are calculated with the Richards-equation (Richards, 1931). The potential evapotranspiration is calculated after Penman-Monteith (Monteith, 1975) and the real evapotranspiration is estimated by using a relation between soil moisture and actual capillary pressure. The calculation of interception is based on a simple bucket approach. Groundwater fluxes are calculated by a two-dimensional flow model which is dynamically coupled to the unsaturated zone. Discharge routing is based on a cinematic wave approach.

The White Volta catchment was subdivided into 7 subcatchments. The outlets of the subcatchments are located at hydrological stations, that simulated discharges can be compared with available measurements. The White Volta flows into Lake Volta, but because of backwater effects which can not be calculated by the model, the outlet of the model setup is not Lake Volta but the station Nawuni. The spatial resolution for this study is  $1 \times 1 \text{ km}^2$  which results in a regular grid of  $411 \times 631$  grid points. The temporal resolution depends on the availability of measurement data during the calibration period which is 24 h for this study. In the vertical the soil is represented by 20 equidistant layers of 1 m thickness each. The model requires digital elevation data, gridded soil properties (derived from the global FAO soil map (FAO, 1971–81)) and gridded land use and hydrogeology information (e.g. saturated hydraulic conductivity, residual and saturation water content) obtained from Martin and van de Giesen (2005). Due to the flat topography in the catchment and an independence of vegetation with altitude the inverse distance weighting method inclusive anisotropy was applied as interpolation method for the observed meteorological input data. The anisotropy impacts significantly the spatial precipitation distribution in the White Volta basin which is characterised by strong latitudinal dependence. Relevant model

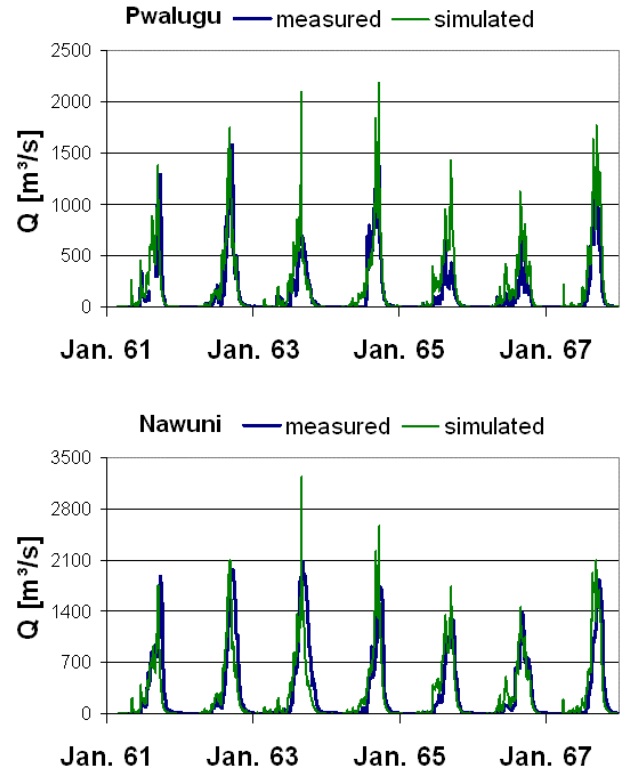


**Fig. 5.** Simulated vs. measured river discharge at Nakong and Nawuni for the calibration period 1968.

parameters were adapted to the specific climatological conditions in the region.

The calibration of the model was done specific for each subcatchment. The recession of direct runoff and interflow both follow a single linear storage concept which required the calibration of the corresponding recession constants. Furthermore, the empirical scaling parameter drainage density ( $d_r$ ) which controls linearly the strength of interflow had to be calibrated. The groundwater model requires the calibration of two empirical parameters, the recession constant ( $k_{rec}$ ), which describes the reduction of the saturated hydraulic conductivity with depth, and the initialisation of the groundwater level in a specific soil layer. The model calibration was performed manually. The calibration period for this study was the year 1968. Especially in areas with weak infrastructure the most important criteria for the selection of a calibration period is the availability of measurement data in the required temporal resolution. For the year 1968 both meteorological and hydrological measurement data were available for both countries Burkina Faso and Ghana. Furthermore, the flow regime of the White Volta was comparatively natural compared to now with the Bagré dam which splits the basin since 1993.

Figure 5 shows the results of the calibration for Nakong (a source basin) and Nawuni (complete basin). It can be seen that the general discharge hydrograph could be simulated

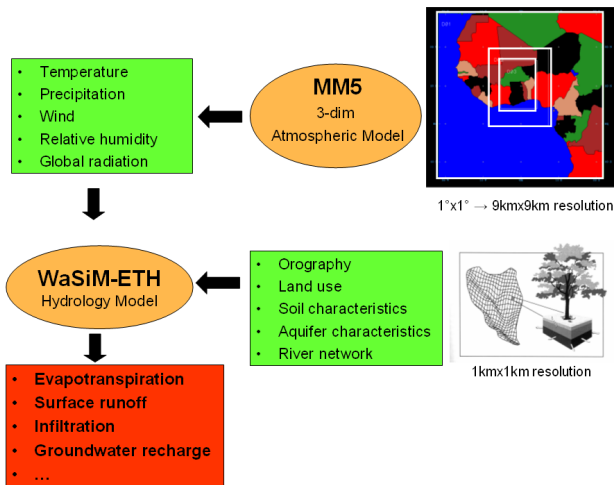


**Fig. 6.** Simulated vs. measured river discharge at Pwalugu and Nawuni for the validation period 1961–1967.

fairly well with Nash-Sutcliffe model efficiencies of 0.72 for Nakong and 0.69 for Nawuni for daily discharges between April and December 1968. Nash-Sutcliffe model efficiencies for all available subcatchments vary between 0.18 and 0.72. The validation of the hydrological model was performed for the period 1961–1967. For this period observation data at Pwalugu and Nawuni were only available. The simulation quality of the validation period is comparable, even slightly better with model efficiencies of 0.74 for Nawuni and 0.64 for Pwalugu (calibration period: 0.48) with one year pre-run (Fig. 6). One reason for the better model efficiencies for the validation period is the relative good agreement between simulation and observation at low flow conditions during the dry season from November until April.

## 5 Coupled meteorological-hydrological modeling

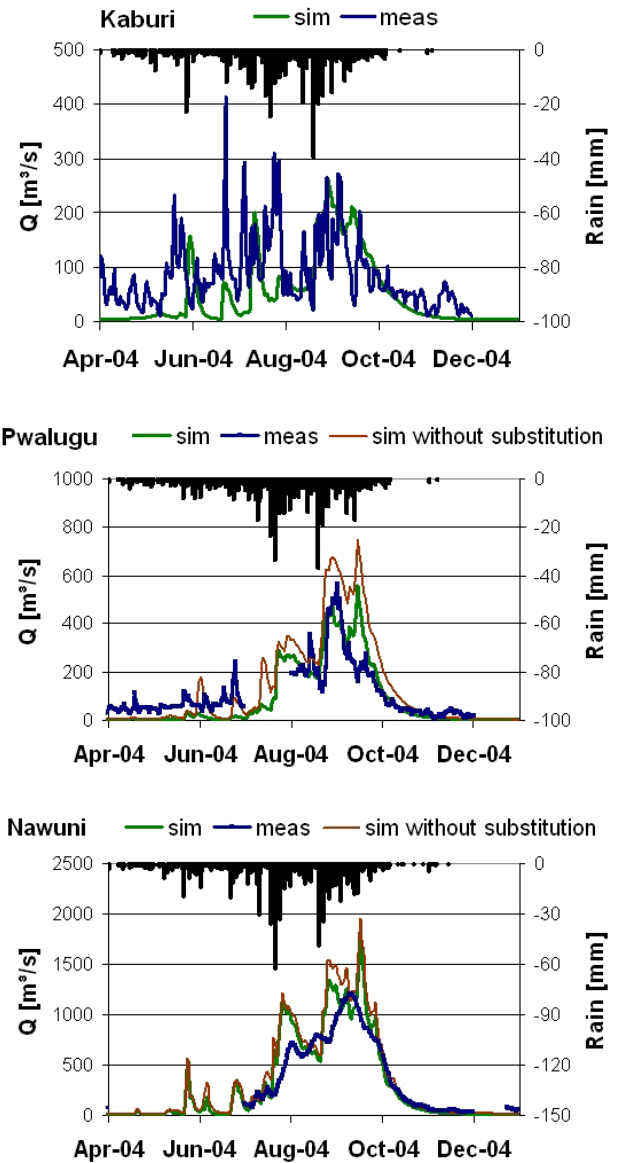
After both models were adapted to the White Volta basin the next step was the one-way coupling of the meteorological and hydrological model (Fig. 7). For the coupling, an interface was developed which generates the required meteorological input data for the hydrological model from the results of the meteorological model. Five meteorological fields are passed: 2m-temperature, precipitation, 2 m-relative humidity, global radiation and 10 m-wind velocity.



**Fig. 7.** Flowchart of the coupled meteorological-hydrological modeling system.

The results from the meteorological simulations for 2004 were processed with the developed interface. For the hydrological modeling the setup from the calibration period 1968 with meteorological observations was taken without recalibration for the coupled simulations. Figure 8 shows the discharge hydrographs of the coupled simulations for three stations along the White Volta. The observed discharges in Kaburi could not be simulated accurately because of the Bagré dam a few kilometres upstream which disturbed the natural flow regime. The discharge hydrographs downstream depend strongly on the management of the Bagré dam. Therefore, the simulated runoff was replaced by the measured runoff at the next gauging station downstream which is Kaburi, to avoid the transmission of errors to the downstream catchments. The discharge hydrographs of the downstream catchments (Fig. 8, Pwalugu and Nawuni) could be simulated fairly well with Nash-Sutcliffe model efficiencies of 0.52 for Pwalugu and 0.63 for Nawuni for daily discharges between April and December 2004. Additionally, the simulated hydrographs without substitution of simulated by measured runoff at Kaburi are plotted to show the influence but also to demonstrate that the model efficiency downstream is not dominated by the substitution which could compensate the results of the coupled meteorological-hydrological model system.

The simulation in Pwalugu underestimates the hydrograph during the first months. The reason is that the storages are empty at the beginning of the model run and these storages have to be filled first before runoff accumulates. So far only one year was simulated but using the same year as pre-run did not work, because the storage processes are long-term (> 1a) in this region. As 2005 data become available the simulations will be continued.



**Fig. 8.** Coupled simulated vs. measured river discharge, plus simulated hydrograph without substitution of simulated by measured values at Kaburi, for Kaburi, Pwalugu and Nawuni.

The coupled meteorological-hydrological simulations show similar model performances as the simulations driven by observations.

## 6 Summary and conclusions

The distributed hydrological model WaSiM-ETH was adapted to the White Volta catchment. The results of the hydrological simulations driven by measuring data for the required meteorological input show that for the calibration and validation period the overall discharge hydrograph could be simulated satisfactorily considering the large catchment



size and weak data availability. The meteorological simulation tends toward overestimating measured precipitation sums especially during rain intensive months. This issue will be analysed further as recent data of 2005 and 2006 and more observations become available. In spite of this trend the results of the coupled meteorological-hydrological simulations show that the general course of the hydrographs can be reproduced fairly well. The discharge is still partly overestimated but in a more moderate magnitude. This means that shortcomings of the meteorological model in simulating the exact location and magnitude of precipitation are partly lessened in sub- and catchment scale considering the results of the hydrological simulations. The model efficiencies of the simulations driven by observations and by meteorological model output are comparable. It is therefore concluded that the meteorological simulations are able to provide the required meteorological fields for hydrological simulations.

This coupled system has the potential for a model based contemporary estimation of the spatial and temporal changes of water balance variables. The basin-wide near-term estimation will provide crucial information for water resources management for the Volta basin.

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