



# SOC stabilization mechanisms and temperature sensitivity in old terraced soils

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Pengzhi Zhao<sup>1</sup>, Daniel J. Fallub<sup>2</sup>, Sara Cucchiaro<sup>3</sup>, Paolo Tarolli<sup>3</sup>, Lisa Snape<sup>4</sup>, Andreas Lang<sup>4</sup>, Sebastian Doetterl<sup>5</sup>, Antony G. Brown<sup>2,6</sup>, Kristof Van Oost<sup>1</sup>

<sup>1</sup>Georges Lemaît Centre for Earth and Climate Research, ELIC, Université Catholique de Louvain, Louvain-la-Neuve, 1348, Belgium;

<sup>2</sup>Tromso University Museum, UiT The Artic University of Norway, Kvaløyen 30, 9013 Tromsø, Norway;

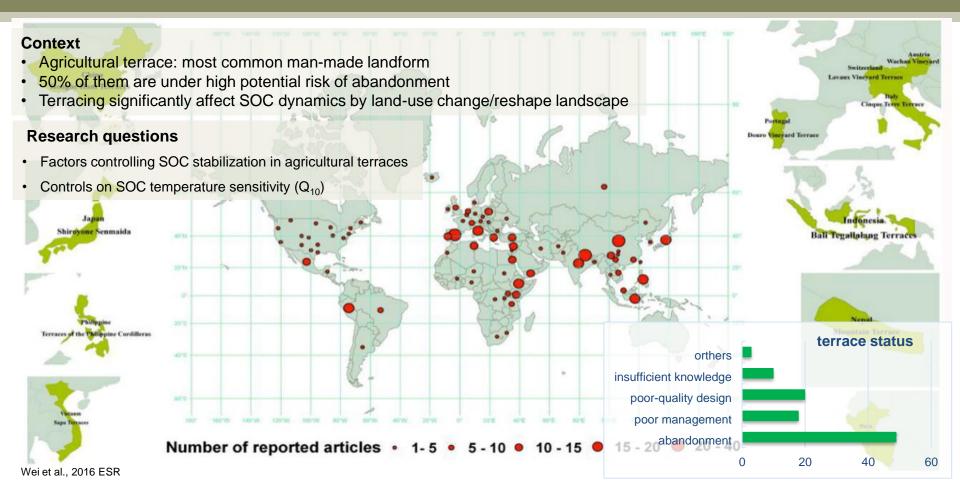
<sup>3</sup> Department of Land, Environment, Agriculture and Forestry, University of Padova, viale dell'Università 16, 35020 Legnaro, Italy;

<sup>4</sup>Department of Geography and Geology, University of Salzburg, Salzburg, 5020, Austria;

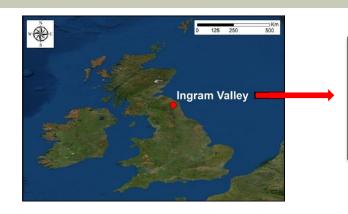
<sup>5</sup>Department of Environmental Systems Science, ETH Zurich, Universitätstrasse 16, 8092 Zürich, Switzerland;

<sup>6</sup>Geography and Environmental Science, University of Southampton, Highfield SO17 1BJ, Southampton, UK

## **Context and research questions**

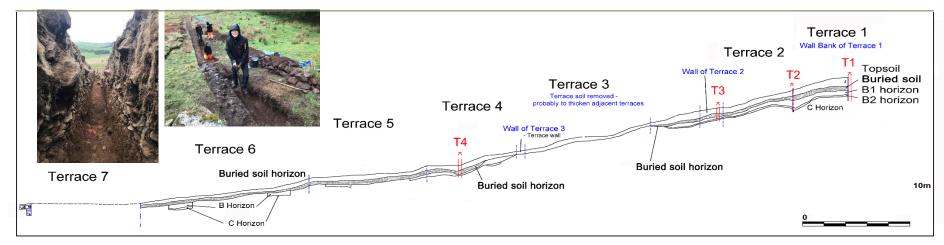


# Studying area

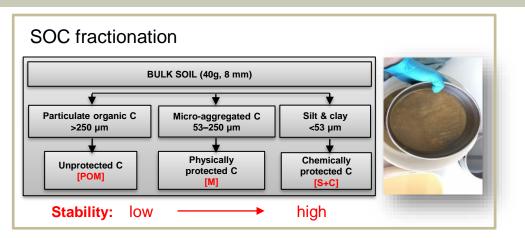


- Northumberland National Park, UK
- > multi-period archaeological landscapes
- > Early Bronze Age c. 1800–1500 BC
- Maritime temperate climate





## Lab analysis



### Elemental composition and pedogenic oxides

- ✓ Rubidium/Strontium (weathering indicator)
- ✓ Sequential pedogenic extractions (Fe, Al, Mn)





#### **SOC respiration -** soil incubation (8 weeks)

30 g 2 mm sieved bulk soil; 350 ml sealed jars; 20 °C and 30 °C; soil respiration + SOC temperature sensitivity ( $Q_{10}$ )

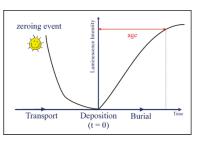




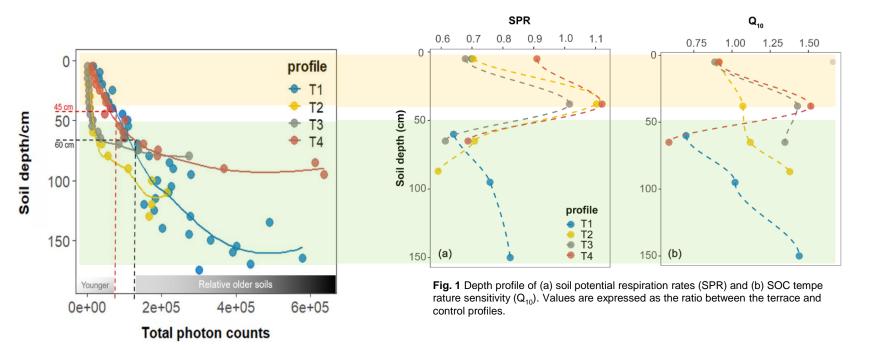
### Soil burial age —— field pOSL

Optically stimulated luminescence (OSL)





# Results —— SOC respiration (SPR) and temperature sensitivity (Q<sub>10</sub>)



Overall, SOC from old soil layers have been protected, but they show higher sensitivity to warming

# Results —— Stabilization mechanisms of terracing SOC

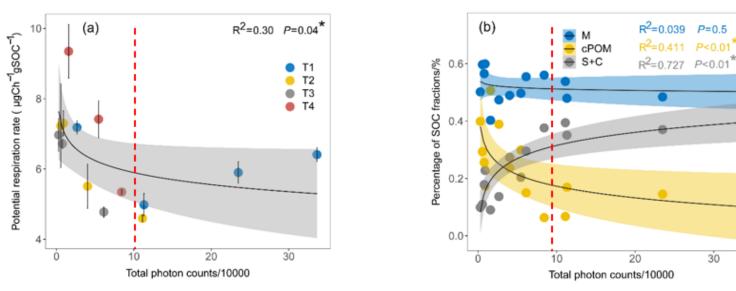


Fig. 2 Relation between terrace soil burial age (total photon counts) and, (a) soil potential respiration rates (SPR) and (b) SOC fractions as a percentage of total SOC (%). Formula: y=log(x).

- older soil horizons (buried layers) tended to have a lower SPR
- The shift to more processed recalcitrant SOC (S+C fracction) with terrace age contributes to SOC stability in terraced soils (Fig. 2b)

### Results —— controls on SOC temperature sensitivity

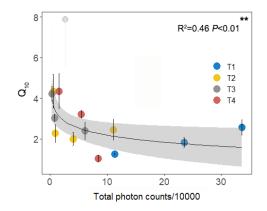


Fig. 3 Relationship between SOC temperature sensitivity to decomposition (Q10) and relative terrace soil burial age (total photo counts).

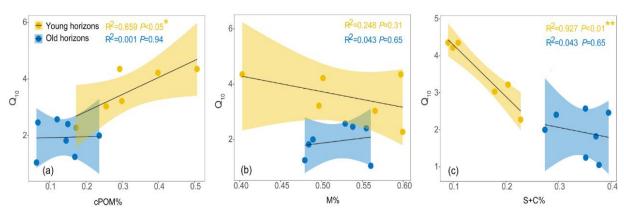


Fig.4 Relationship between SOC temperature sensitivity (Q10) and (a) unprotected SOC (cPOM%), (b) physical protected SOC (M%) and (c) mineral protected SOC (S+C%) for relative younger and older terrace soil horizons, respectively. \*= P<0.05: \*\*P<0.01.

Table 1 Correlation between SOC fractions and pedogenic oxides

Young soil layers											
cPOM m	$Al_p$	Fe <sub>p</sub>	$Mn_p$	$AI_o$	Feo	Mn <sub>o</sub> -0.81	$AI_d$	Fe <sub>d</sub>	Mn <sub>d</sub> -0.76		
s+c	0.78	-0.88		0.89	0.88	0.83	0.84		0.88		
Old soil layers											
cPOM m s+c											

#### Young soil horizons

SOC mineral protection attenuate the SOC intrinsic temperature sensitivity by reducing the availability of SOC substrate to decomposers

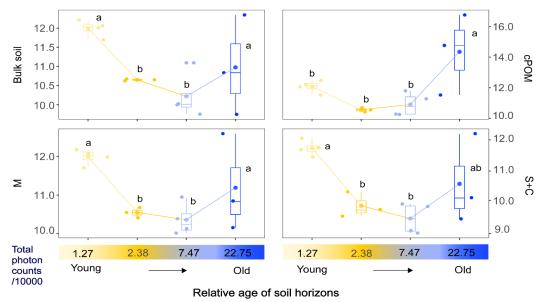
#### Old horizons?

### Results —— controls on SOC temperature sensitivity

**Table 2** Relationship between SOC temperature sensitivity ( $Q_{10}$ ) and C:N ratios of bulk soil and SOC fractions.

_	Bulk soil	cPOM	M	S+C
$Q_{10}$	0.60*	0.03	0.61*	0.62*

<sup>\*</sup> P<0.05. N=13.



**Fig. 5** C:N ratios for bulk soil and SOC fractions along with the gradient of terrace soil burial age (total photon counts). Significant differences in C:N ratios between soil age gradient are indicated by different lowercase letters (*P*<0.05).

#### Old horizons

Higher C:N ratio (lower quality) of SOC lead to a higher temperature sensitivity of SOC stored in buried horizons

### Conclusions

- Soil burial due to terracing provides a C stabilization mechanism.
- With increasing burial age, the SOC pool composition shifts from particulate OC to mineral protected OC pool.
- Both soil C:N ratio (C quality) and SOC mineral protection regulate Q<sub>10</sub>
- The dominant mechanism controlling this temperature sensitivity depends on the burial age



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pengzhi.zhao@uclouvain.be

kristof.vanoost@uclouvain.be

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